

# Hands-on Tutorial on Phonon calculations

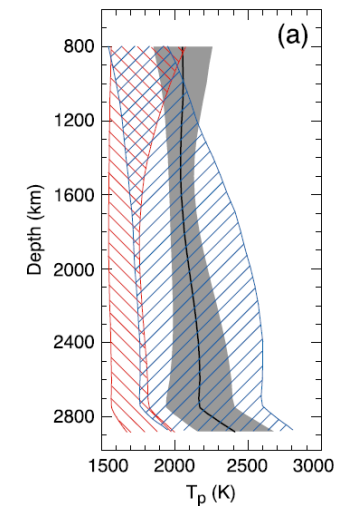
- > Adams DJ  
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- > Overview
  - Introduction: How include thermal effects in theory?
  - Phonons, harmonic approximation: Theory
    - Harmonic oscillator
    - Thermodynamics (Quasi Harmonic Approximation, QHA)
    - Contributions from long wavelength phonons
  - Structure of the Earth's mantle: Applications QHA
  - Extension of Quasiharmonic approximation: Decoupled anharmonic Mode approximation (DAMA)
  - Conclusions and outlook

# Introduction: How include thermal effects in theory?

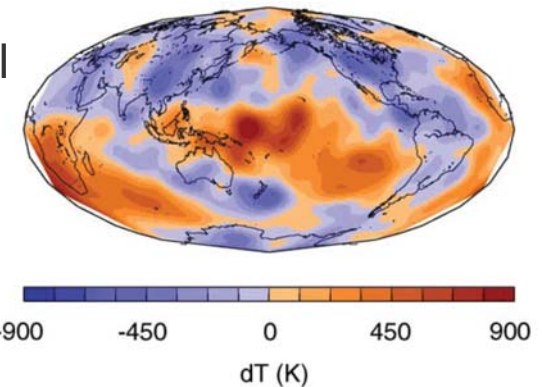
- > Theoretical calculations often performed at  $T=0$  K !
- > In Earth interior minerals are commonly at  $T>0$  K
- > Measured wave velocities

$$d \ln V_p = \frac{\partial \ln V_p}{\partial T} dT + \frac{\partial \ln V_p}{\partial C} dC + \frac{\partial \ln V_p}{\partial F} dF$$

$$d \ln V_s = \frac{\partial \ln V_s}{\partial T} dT + \frac{\partial \ln V_s}{\partial C} dC + \frac{\partial \ln V_s}{\partial F} dF$$



- > Numerical inversion  $\implies$ 
  - composition ( $C$ ), temperature ( $T$ ), fraction of partial melt ( $F$ )



Masters, G. et. al. (2000). In Karato, S. et al. Deep Interior: Mineral Physics and Tomography from the Atomic to the Global Scale, Washington, Am. Geophys. Union.

Deschamps, F. and Trampert, J. (2003). Phys. Earth Planet. Int., 140:277–291

Deschamps F, Trampert J / Earth and Planetary Science Letters **222** (2004)

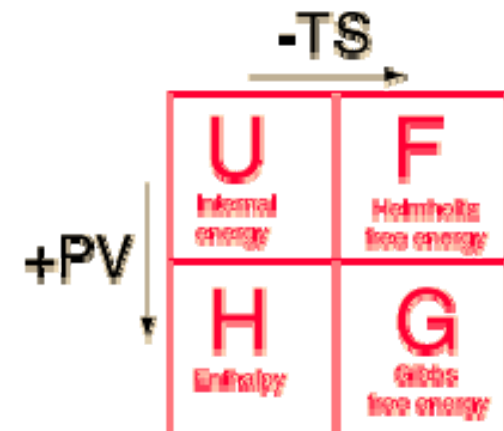
161–175

# Background Thermodynamics

- > Free energy  $G(T,P)$  is needed to calculate phase equilibria
- >  $G(T,P)$  can be obtained from  $F(T,V)$

$$G(P,T) = F(V,T) + P \cdot T = \underbrace{E_{0K}(V)}_{=U} + E_{ZP}(V) - T \cdot S(T,V) + P \cdot V$$

- >  $E_{0K}(V)$  can be obtained from static calculations (geometry optimization)
- >  $E_{ZP}(V)$  is a (small) correction, taking into account the quantum nature of the atomic cores
- >  $E_{ZP}(V) - T \cdot S(T,V)$  can be calculated from partition function



# Background Harmonic Oscillator

> Move 1 atom along 1 degree of freedom (N atoms, 3N-3 degrees of freedom)

> Curvature of potential energy surface determines *all* vibrational frequencies (e.g. Wu, 2008)

> (Quasi) Harmonic approximation => atoms are “connected by springs” and experience the potential

$$U = \frac{k}{2} (q_\mu)^2$$

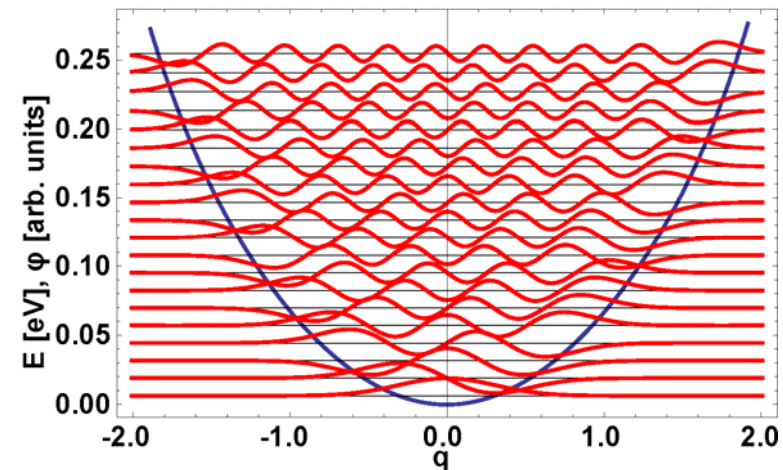
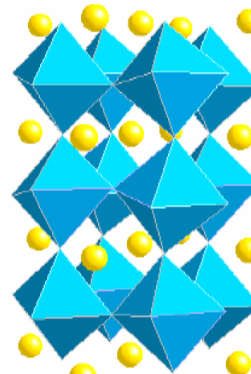
> Energies for harmonic potential are equidistant

$$\mu \in [1, 3N - 3]$$

> Indices:

- Accounts for degree of freedom
- J accounts for energy level

> Theory breaks down for negative curvatures of potential energy surface



$$\omega_\mu = \sqrt{\frac{1}{m_\mu} \left. \frac{\partial^2 V(q_\mu)}{\partial q_\mu^2} \right|_{q_\mu=0}}$$

$$\epsilon_\mu^j = \hbar \cdot \omega_\mu \cdot \left(\frac{1}{2} + j\right) \text{ and } j \in \mathbb{N}_0$$

# Background Thermodynamics

- > The partition function  $Z$  gives access to all thermodynamic quantities
- > The free energy is the thermodynamically relevant potential at  $T > 0$
- > Thermodynamic energy
- > Heat capacity
- > Elastic constant tensor

$$Z = \sum_i e^{-\frac{\epsilon_i}{k_B T}}$$

$$A = -k_B T \ln Z$$

$$\langle E \rangle = k_B \cdot T^2 \frac{\partial \ln Z}{\partial T}$$

$$c_V = \frac{\partial \langle E \rangle}{\partial T}$$

$$c_{ij}^T(T, P) = \left[ \frac{\partial^2 A}{\partial \epsilon_i \partial \epsilon_j} \right]_P$$

$$G(P, T) = F(V, T) + P \cdot T = \underbrace{E_{0K}(V)}_{=U} + \underbrace{E_{ZP}(V) - T \cdot S(T, V)}_A + P \cdot V$$

# Background Phonons

- > In solid, the lowest energy excitations are collective vibrations of atoms: phonons

$$\Phi = \begin{bmatrix} \frac{d^2 E}{dq_1 dq_1} & \dots \\ \dots & \frac{d^2 E}{dq_\mu dq_\nu} \end{bmatrix}$$

- > They can be obtained from force constant matrix

$$-\frac{dE}{dq_\mu} = F_\mu \Rightarrow \frac{d^2 E}{dq_\mu dq_\mu} = -\frac{dF_\nu}{dq_\mu}$$

- > The eigenvalues of the dynamical matrix  $D$  are eigenenergies

$$D = \begin{bmatrix} \frac{1}{\sqrt{m_1 \cdot m_1}} \frac{d^2 E}{dq_1 dq_1} & \dots \\ \dots & \frac{1}{\sqrt{m_\mu \cdot m_\nu}} \frac{d^2 E}{dq_\mu dq_\nu} \end{bmatrix}$$

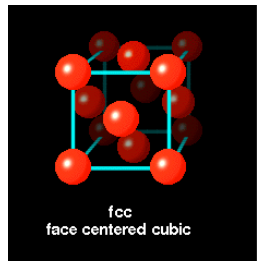
- > The eigenvectors of the dynamical matrix  $D$  are the polarization vectors

# Example Dynamical matrix

> Dynamical matrix

In cartesian coordinates

$$D = -\frac{1}{\sqrt{m_\mu \cdot m_\nu}} \frac{dF_\nu}{dq_\mu} = \begin{pmatrix} -6. & 0. & 0. & 0.2 & 0.4 & 0. \\ 0. & -6. & 0. & 0.4 & 0.2 & 0. \\ 0. & 0. & -6. & 0. & 0. & 0. \\ 0.2 & 0.4 & 0. & -6. & 0. & 0. \\ 0.4 & 0.2 & 0. & 0. & -6. & 0. \\ 0. & 0. & 0. & 0. & 0. & -6. \end{pmatrix}$$



> Polarization vector  
— Collective modes

$$\varepsilon = \begin{bmatrix} 0.5 \\ 0.5 \\ 0 \\ -0.5 \\ -0.5 \\ 0 \end{bmatrix}$$

$$\varepsilon_i = \begin{pmatrix} 0.5 & 0.5 & 0 & 0 & 0.5 & 0.5 \\ 0.5 & -0.5 & 0 & 0 & -0.5 & 0.5 \\ 0 & 0 & 1. & 0 & 0 & 0 \\ -0.5 & 0.5 & 0 & 0 & -0.5 & 0.5 \\ -0.5 & -0.5 & 0 & 0 & 0.5 & 0.5 \\ 0 & 0 & 0 & 1. & 0 & 0 \end{pmatrix}$$

> Eigenvalue

$$D = \begin{pmatrix} 6.6 & 0. & 0. & 0. & 0. & 0. \\ 0. & 6.2 & 0. & 0. & 0. & 0. \\ 0. & 0. & 6. & 0. & 0. & 0. \\ 0. & 0. & 0. & 6. & 0. & 0. \\ 0. & 0. & 0. & 0. & 5.8 & 0. \\ 0. & 0. & 0. & 0. & 0. & 5.4 \end{pmatrix}$$

## Background

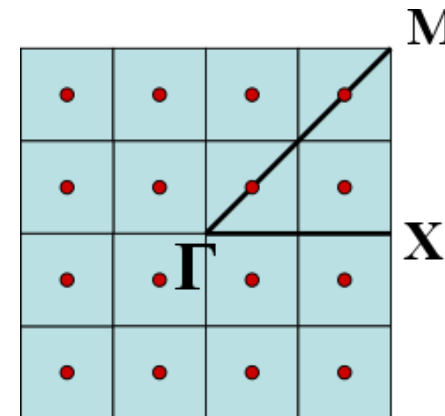
### Contribution of long wave-length phonons

- > Free energy of solid

$$A = U + \underbrace{\sum \frac{1}{2} \hbar \omega_{\mu} + k_B T \ln \left( 1 - e^{-\frac{\hbar \omega_{\mu}}{k_B T}} \right)}_{f(\omega_{\mu})}$$

- > Contribution from long wave-length phonons are important

- > Avoid large supercells through summation over Brillouin Zone
  - In small cell determine force constants
  - Recompute Dynamical matrix at several long wave-length phonons
  - Sum up contributions using multiplicity  $w_{\nu}$

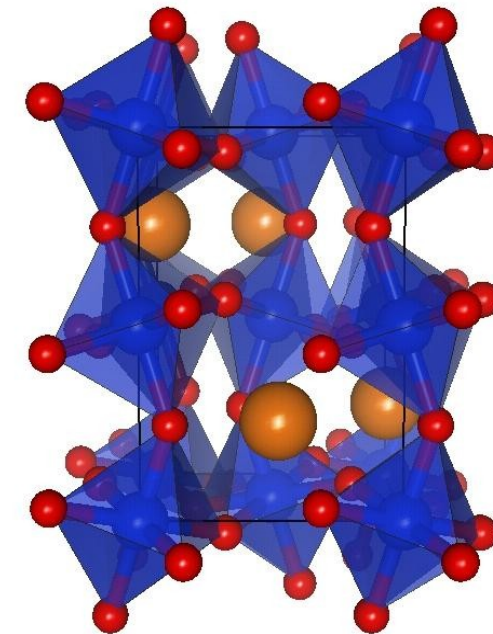
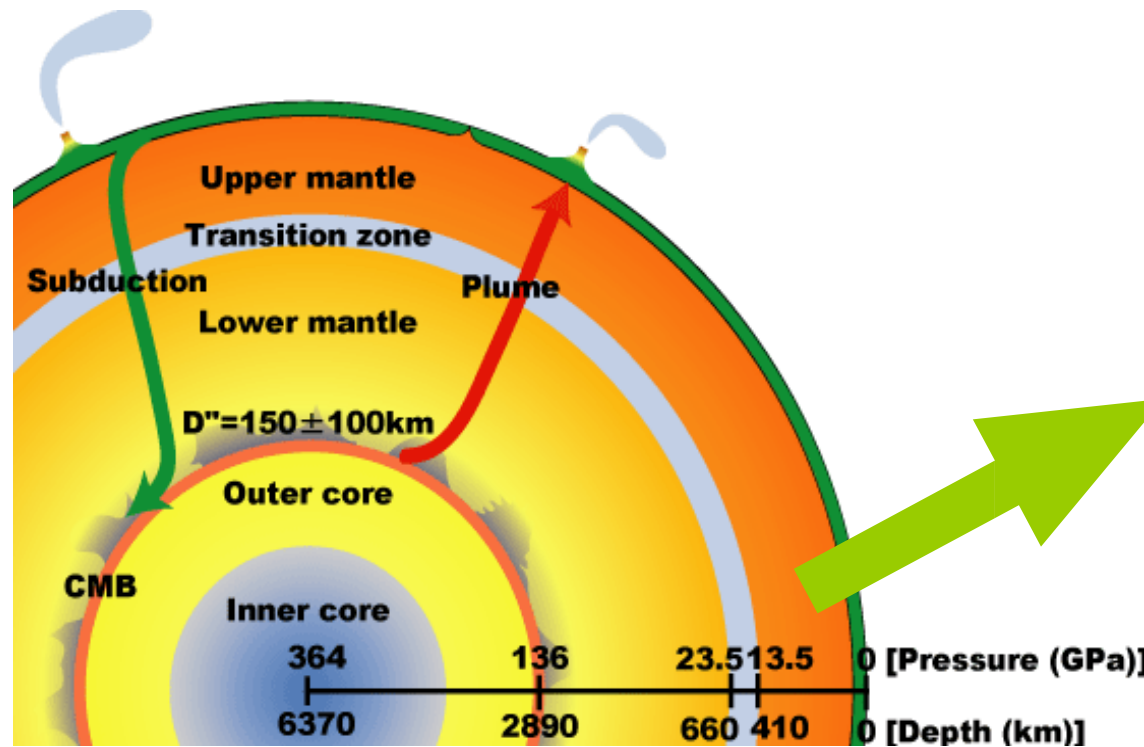


$$A \rightarrow U + \sum_{\nu << 3N-3} w_{\nu} \cdot f(\omega_{\nu})$$

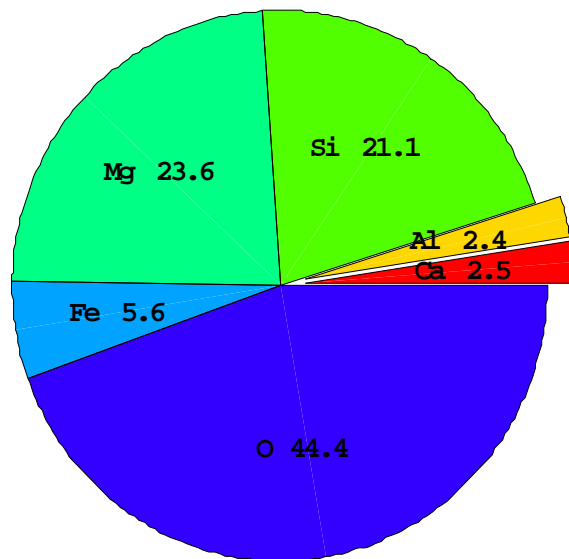


# Composition of the Earth's (lower) mantle

- >  $\text{MgSiO}_3$ : most abundant constituent in the Earth's lower mantle
- > Orthorhombic distorted perovskite structure ( $Pbnm$ )
- > Its stability is important for understanding deep mantle (D'' layer)



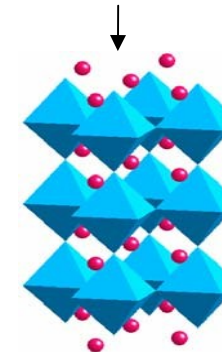
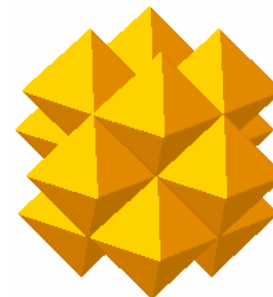
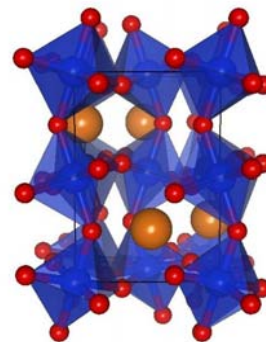
# Lower mantle composition



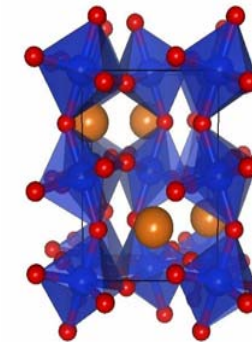
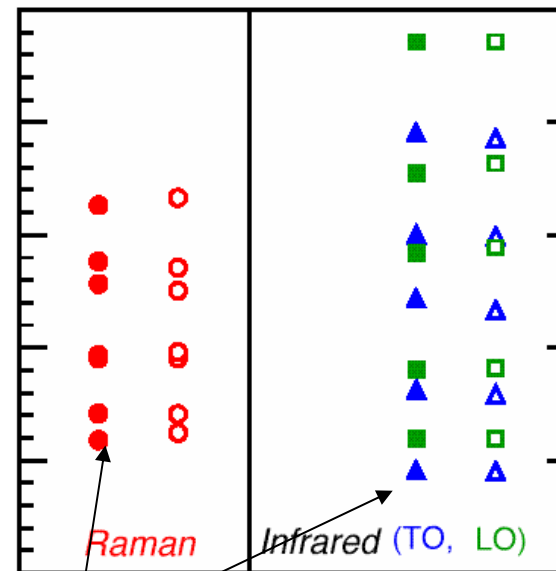
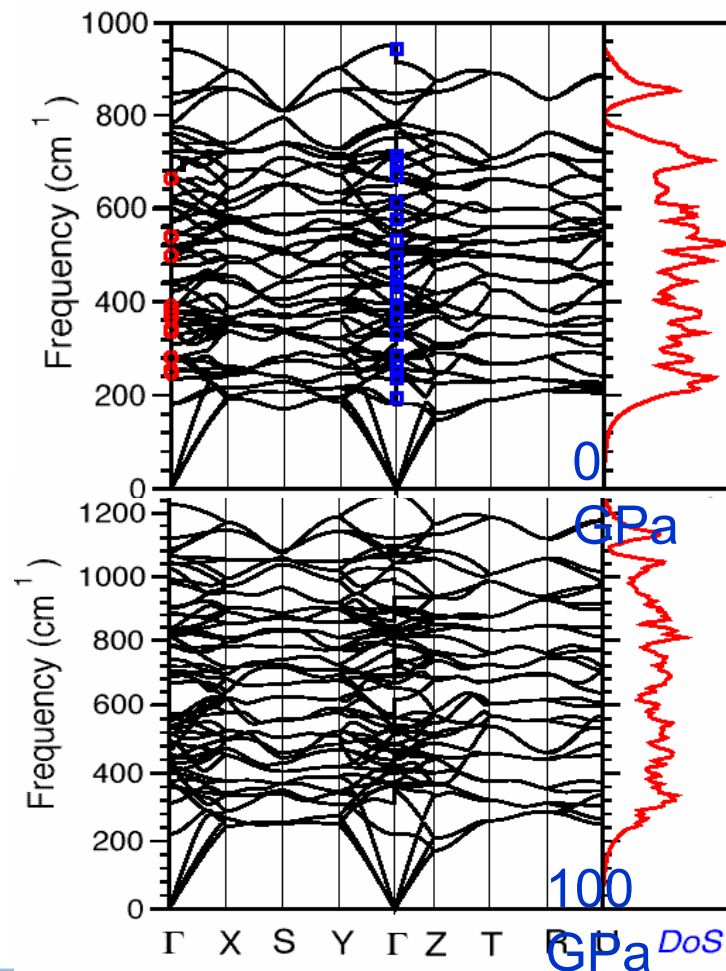
## Lower mantle:

- > 53 vol.% of the Earth
- > 75 wt.% -  $\text{MgSiO}_3$  (post)-perovskite
- > 18 wt.% -  $(\text{Mg,Fe})\text{O}$  magnesiowüstite
- > 7 wt.% -  $\text{CaSiO}_3$  perovskite

Not suitable for harmonic approximation



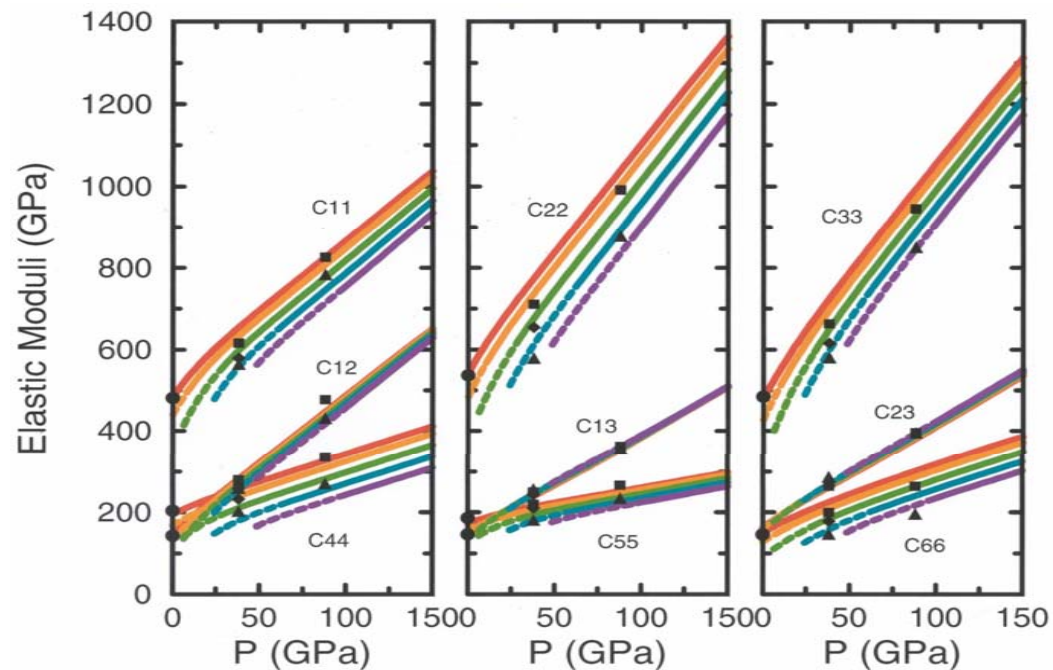
# Phonon dispersion of MgSiO<sub>3</sub> perovskite



Theory: Karki, Wentzcovitch, de Gironcoli, Baroni (2000) PRB 62, 14750  
 Exp: Raman (Durben and Wolf, 1992)  
 Infrared (Lu et al. 1994)

# Temperature dependent elastic constants MgSiO<sub>3</sub> perovskite

$$c_{ij}^T(T, P) = \left[ \frac{\partial^2 A}{\partial \epsilon_i \partial \epsilon_j} \right]_P$$



Wentzcovitch, Karki, Cococciono, de Gironcoli (2004) Phys. Rev. Lett.

# Hands-on Tutorial on Phonon calculations

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## > Conclusions

- Harmonic theory treats vibrations as if they did not interact
- System is equivalent to a collection of independent harmonic oscillators
- Energies used to compute partition function  $Z$  and the free energy,  $A(T)$ .
- Through the free energy all properties are accessible
- DAMA allows to calculate the free energy even for materials with dynamic instabilities
- Dynamic instabilities are common in high temperature phases
- Good agreement between QHA (and DAMA) and experiment, e.g. for phase transition cryolite

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